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"Infrasonic monitoring of snow avalanche activity:

what do we know? where do we go from here?"

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Abstract

The first recordings of acoustic signals associated with avalanches go back to the 1970's when premonitory low frequency acoustic emission was observed in the snow cover just before the setting off of natural avalanches (Gubler, Sommerfeld, 1977). Recently, it has been demonstrated that avalanches also produce strong infrasonic vibrations in air during their movement (Bedard, 1986, Hejda, 1995). These infrasonic vibrations propagate great distances and can follow the natural relief. It was then a question of automatically detecting this emission to monitor avalanche activity, with a view of improving forecasting and assisting in the study of certain aspects of avalanche dynamics.

With the double aim of pursuing investigations on the mechanics of the acoustic emission from avalanches and designing an automatic detection system, our laboratory developed specific instrumentation, the essential link of which is an array of infrasonic microphones associated with a goniometer. Systematic measurements of the infrasonic emissions are being carried out, not only of the natural or released avalanches themselves, but of all events with the aim of characterizing the former in relation to « infrasonic noise ».

Keywords: avalanches, infrasounds, acoustic sensors.

Introduction

Recent work demonstrated that snow avalanches have an acoustic emission, particularly strong below 20 Hz, within the infrasonic range (Chritin V., Rossi M. 1995). These emissions can be detected by an appropriate system including infrasound microphones and signal processing in real time. Thus, one can hope to automatically record avalanche activity over a mountain range: such information could be of prime importance to forecasting systems (Chritin V. and *al* 1996). Since 1995, our laboratory has at its disposal an experimental system, made up mainly of an infrasound goniometer, allowing the localisation of an infrasound source - azimuth and elevation are computed in real time - and an acoustical analyser - performing a frequency analysis. These data are processed by automatic recognition procedures to answer the question of whether the detected acoustical event is a snow avalanche or not? One of the most difficult engineering problems encountered was the design of the sensor system and its set-up in very harsh mountainous conditions. This paper presents the assigned system of electroacoustic transducers designed by our laboratory and its performance relating to detection of infrasonic events such as snow avalanches.

The experimental set-up

Our experimental set-up is installed at Anzère ski resort at an altitude of 2400m (Fig. 1). This site is particularly well suited because numerous avalanches are set off there as a precautionary measure - thus allowing their acoustic characteristics to be studied. Furthermore, it is very close to the experimental site of "La Sionne" for the study of avalanche dynamics. The processing electronics are installed in a shelter, shielded from the bad weather, but the microphones have to be located outdoors, in what is obviously a very hostile environment.

Fig. 1

One of the main difficulties was in fact related to the system of acoustical sensors. Questions of particular concern included, how to protect them, where to install them, how to establish the best conditions for infrasound detection? Furthermore, what are the most suitable microphones?

The first set-up (Rossi M., Chritin V. 1995) used four outdoor electrostatic measurement microphones mounted on masts and protected by large shields (Fig. 2). They were in the same horizontal plane and formed a half-cross (Fig. 3).

This set-up had the following disadvantages:

- microphones on masts are exposed to extreme weather conditions: wind, rain, snow, thunder, risk of ice forming, intense sun radiation;
- their masts constitute obstacles on the ski slopes;

- setting-up of a mast with a concrete base represents substantial additional cost and environmental impact;
- wind induces noise from the turbulence on the obstacle formed by the microphone and its shield, which adds to the noise from the wind on other obstacles; the system is "blinded" by an insignificant amount of wind and its range is thus reduced;
- the performance of an outdoor electrostatic measurement microphone is not very well matched for infrasound detection: the low cut-off frequency is too high, the upper one is unnecessarily high, the noise is too high (it follows a $1/f$ law); a costly high input impedance preamplifier is required.

Fig. 2

Design of a dedicated microphone

In 1995 we designed a dedicated microphone for infrasound detection, called ECHO, suitable to be placed in a pit under the snow cover, its acoustical input port protected by a shield. Four ECHOs were installed in November 1995 and are still working perfectly well today.

This microphone was designed to meet the following criteria: low frequency (down to 1.6 Hz), low noise (improvement of at least 20 dB of the signal to noise ratio), low cost, simple electronics (low cost), great robustness, low sensitivity to the environment, bulk and size not critical, but simple installation without excessive extra costs.
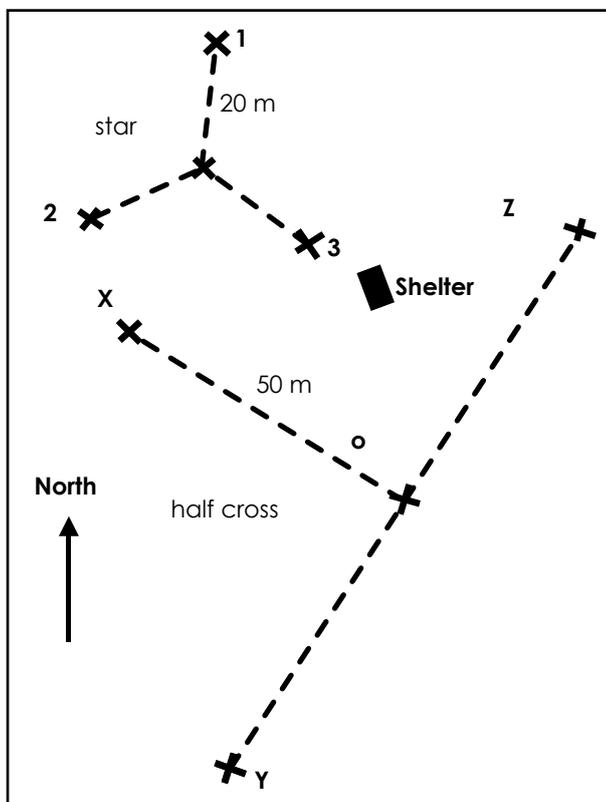
Fig. 3

An 18 month investigation was carried out to test these new microphones and to assess their performance. The second experimental set-up was a star (Fig. 3) of four ECHOs in pits under the snow cover, with four supplementary outdoor electrostatic measurement microphones placed under the shields of the ECHO (Figs. 4 and 5). One of these microphones was a low noise type, not intended for outdoor use in an hostile environment,

so it was out of order after a few months. The first microphone set-up was maintained in use for comparison purposes.
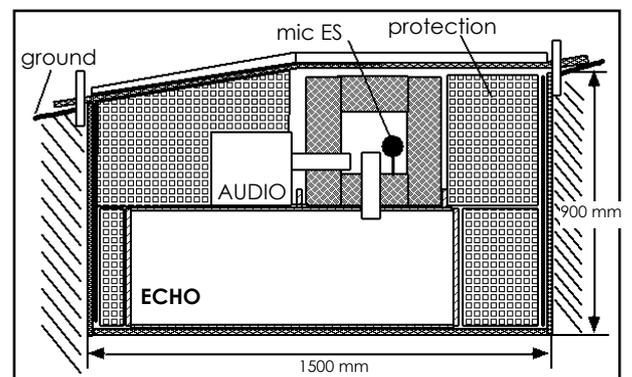
Fig. 4

Fig. 5

Sensitivity to wind noise

Figure 6 shows the outputs of one electrostatic measurement microphone (abbreviation ES) on a mast and of one ECHO beneath the snow cover, along with the wind speed over ground. Clearly, the ECHO performs better, its wind induced noise is lower than that of the ES.
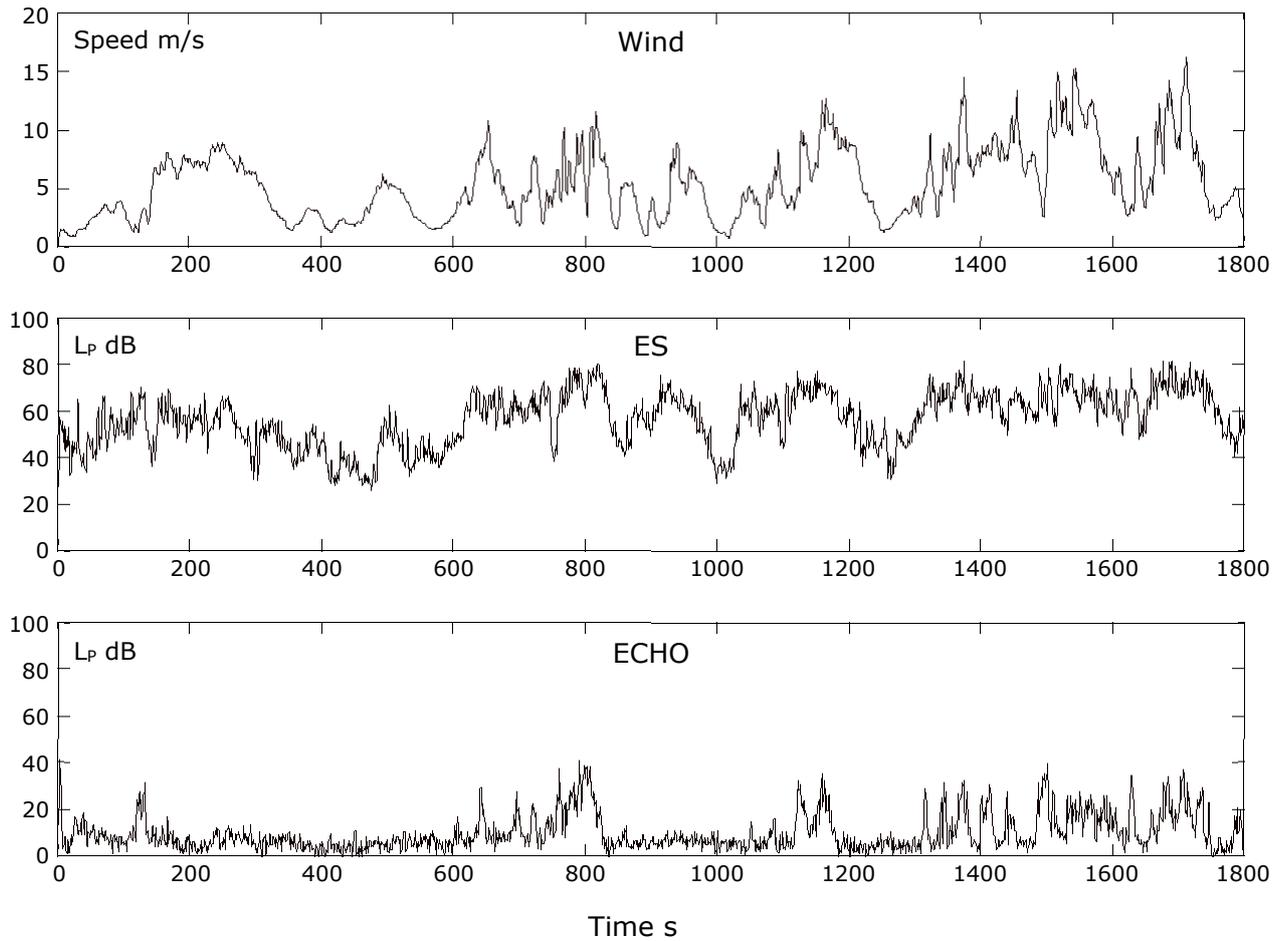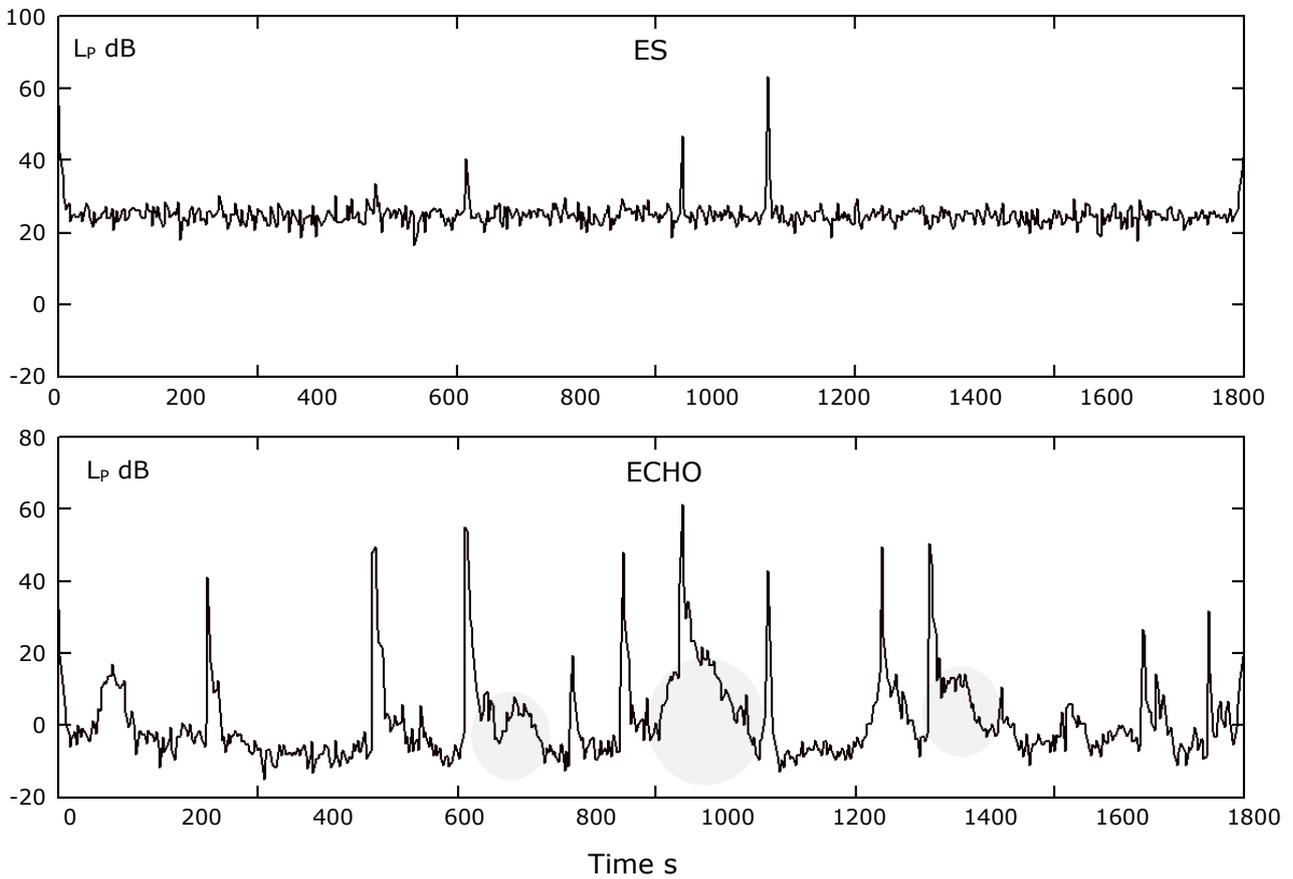
Fig. 6

Fig. 5

Microphone noise

Figure 7 shows the sound events during a phase of precautionary release of avalanches using explosives dropped from helicopters, as seen by one electrostatic measurement microphone ES and one ECHO both located in a pit as shown in Fig. 4. The output of the ES is very noisy (upper graph), however the acoustical events are clearly detected by the ECHO (lower graph), particularly the snow slides occurring after the explosions (they were observed visually).

Environment infrasonic noise

Figure 8 exhibits the typical ambient noise levels versus time (four days) as measured by one low noise electrostatic measurement microphone (ES) and one ECHO. ES microphone is lying on the ground (no snow cover) under a wind shield and the acoustical input port of the ECHO is just under ground level. An analysis in third octave bands was performed from 1.6 Hz to 16 Hz. A mean value of the wind speed was simultaneously recorded and is represented along the octave band levels. As shown in Figure 9, in case of a light wind the ambient noise is practically the same for the two microphones, but in strong wind, the noise is less in the lower part (from 2 to 5 Hz) for the ECHO than for the ES.
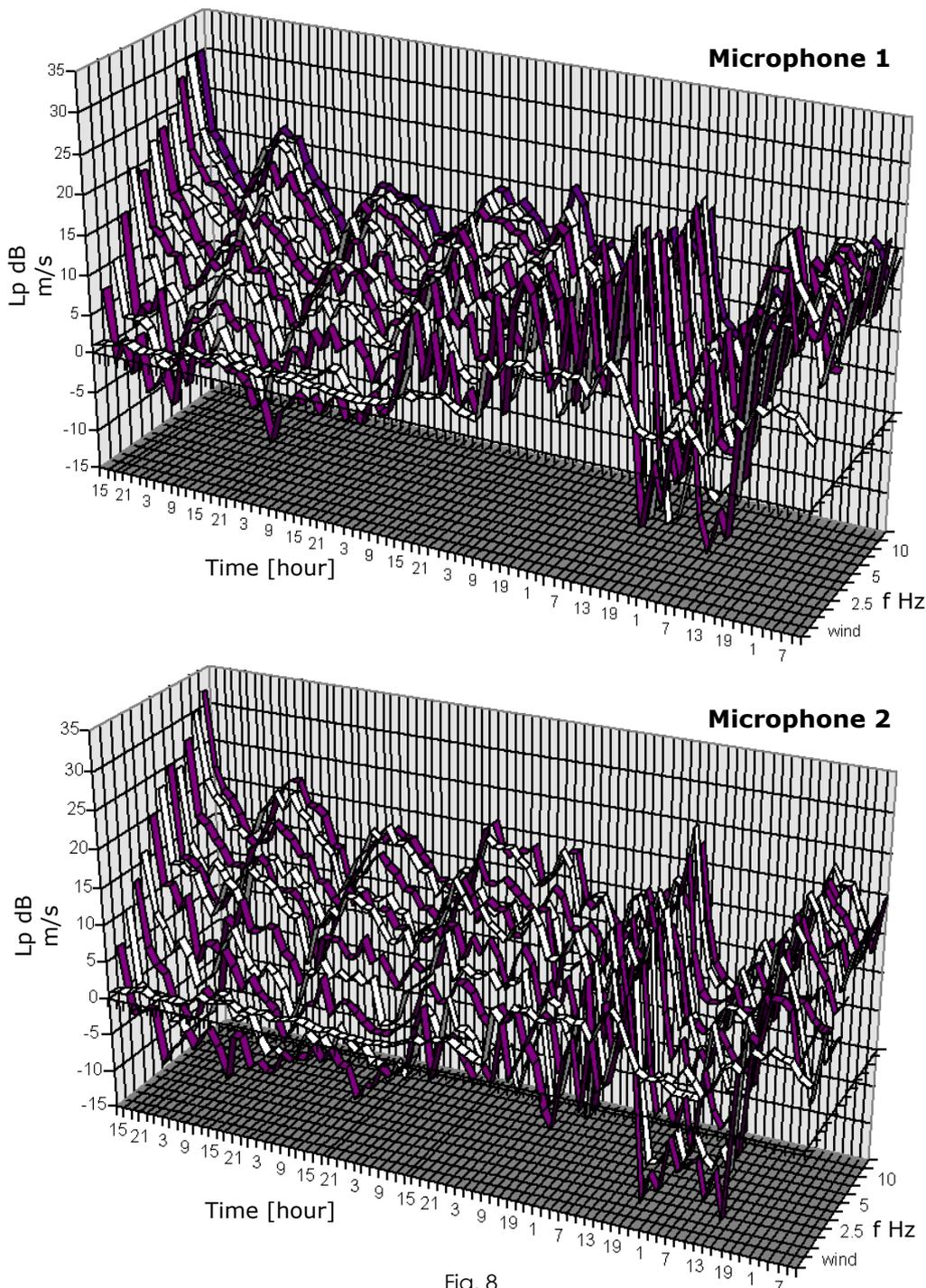

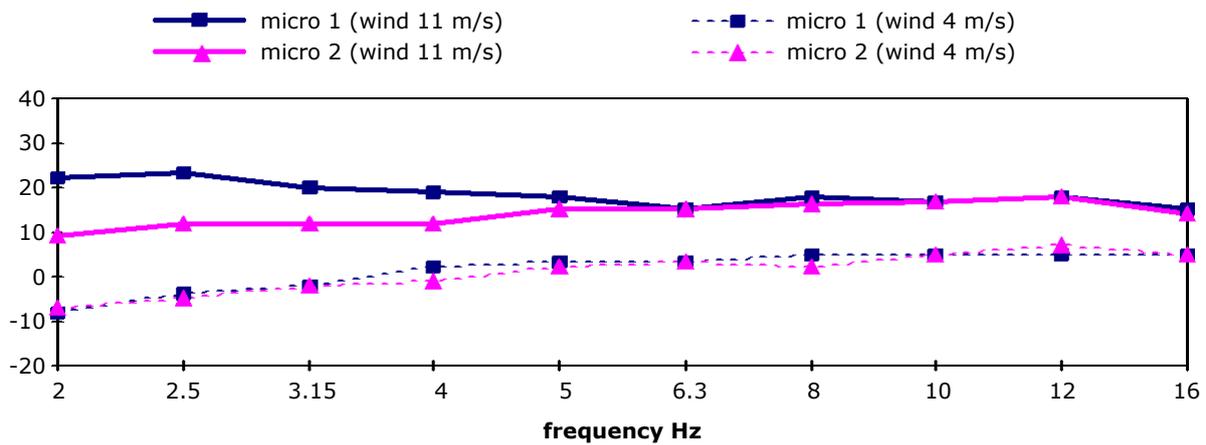**Microphone 1**

Microphone 2

Fig. 8

Fig. 9

Snow cover

Systematic experiments were carried out in order to determine the effect of the snow cover on the acoustical sensors. For this purpose, two ECHOs, were installed as shown by in Figure 10. The snow cover varied according to the snowfalls. The ECHO above the snow cover was taken away between measurement periods. It was sometimes necessary to top up the snow 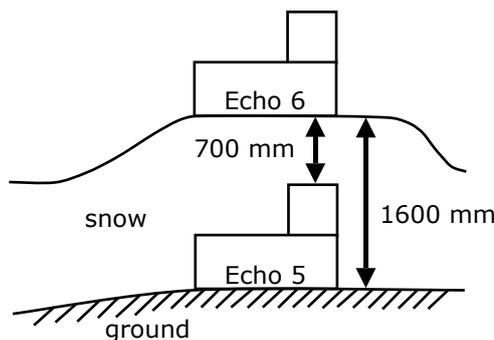cover which tended to get packed down. Because of it, this more or less porous layer could often be very packed down and once even had a layer of ice. Comparison between the signals picked up by both microphones enabled us to determine the effects of the snow cover. Figure 11 exhibits typical waveforms of
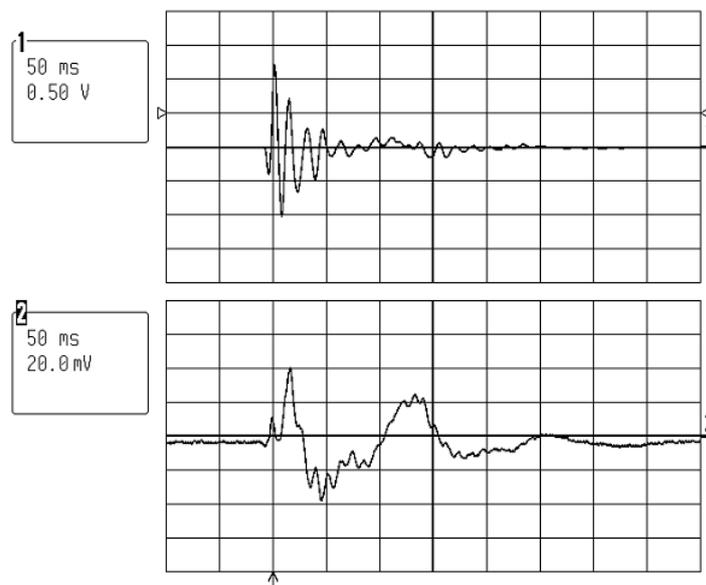

Fig. 10

explosions (lower trace: ECHO below snow cover; upper trace: ECHO above snow cover). Overall 1-100 Hz range, an attenuation by a factor of about 25 is observed, as well as a low-pass filtering effect. Figure 12 shows the correlation between the previous waveforms: a delay of 21 ms for the signal picked up by ECHO below is observed. This value is virtually independent of the distance and direction of location of the explosion (many measurements were made for different locations of the explosions). The snow conditions this year did not allow us to compile sufficient data to correlate this with snow depth.

The observation of acoustic emissions from avalanches led to the same conclusions. Figure 13 shows the waveforms of an avalanche (lower trace: ECHO below; upper trace: ECHO over snow cover) and Figure 14 exhibits their correlation. An estimation of the transfer function between the sound pressures at the acoustical ports of the ECHO above snow cover and below snow cover led to the graph shown in Figure 15. In the infrasonic range, in which we are interested for avalanche monitoring, the observed low-pass filtering and attenuation are not excessive.
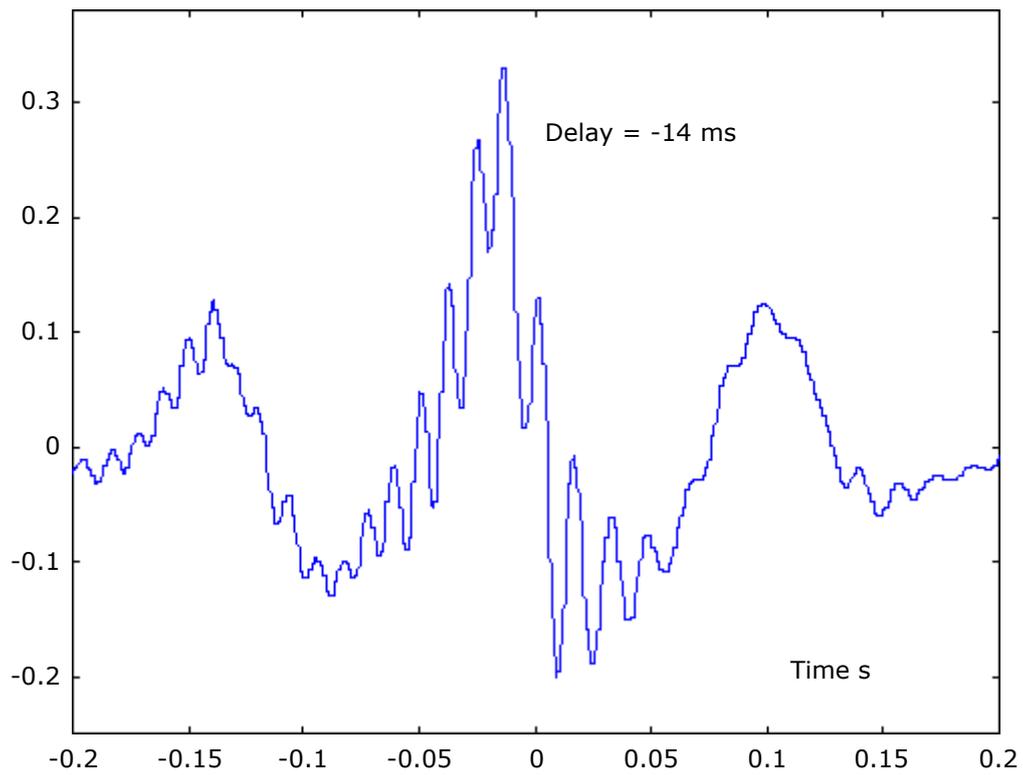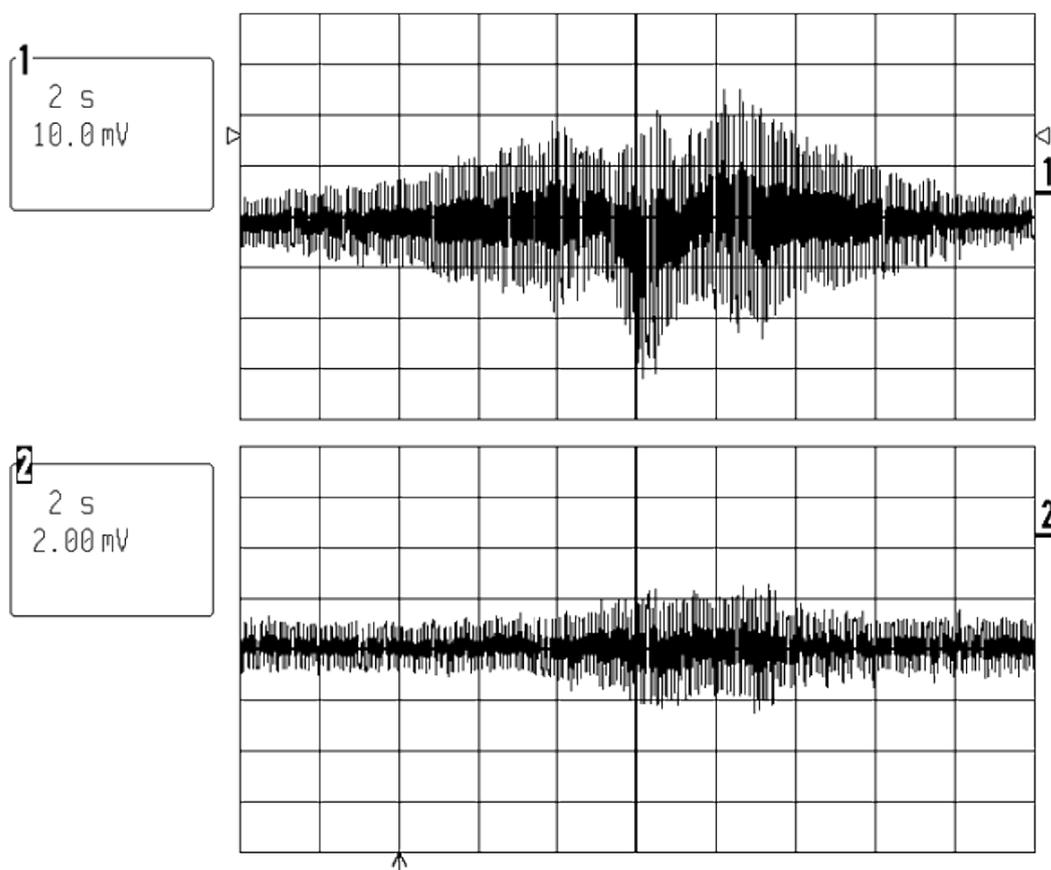

Fig. 11

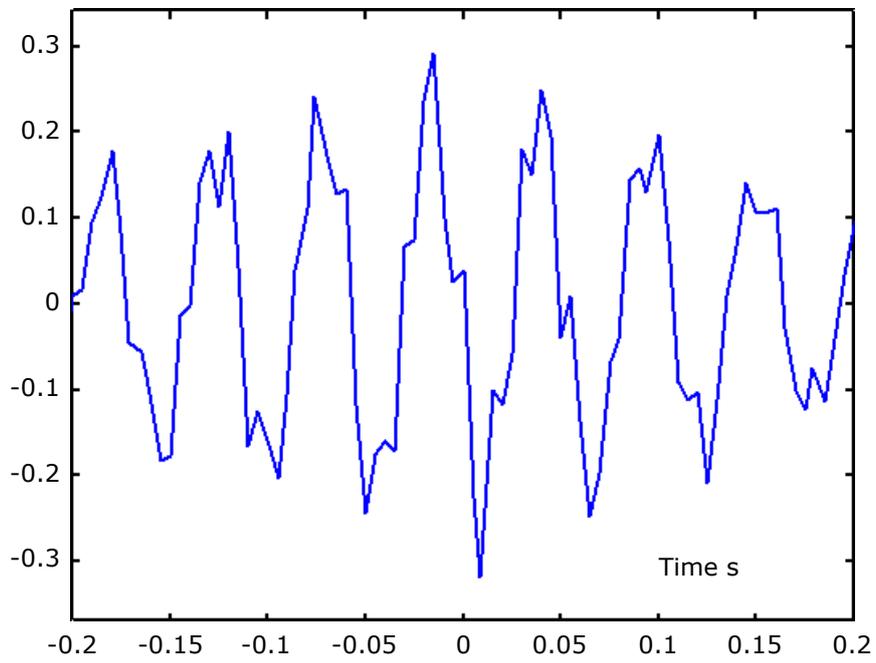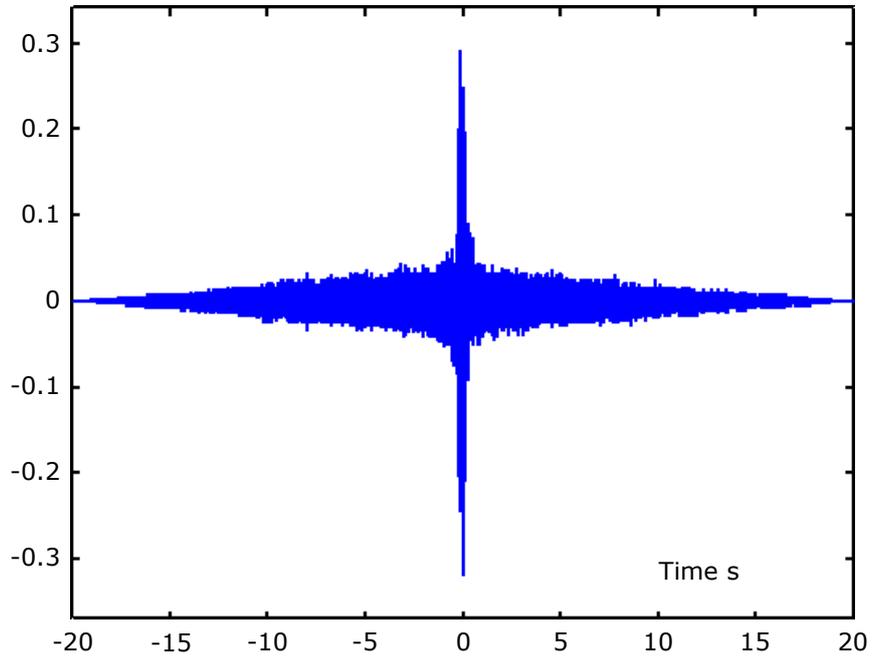Delay = -14 ms

Time s

Fig. 12

Fig. 13

Fig. 14

Conclusion

In conclusion, the ECHO microphone system presents undeniable advantages compared to what was available before and does indeed fulfil the requirements for which it was designed. To illustrate this, Figure 16 shows the waveform and the spectrograh of a succession of avalanches picked up by an ECHO below a snow cover of 2 m.

These results are very encouraging, because the development of avalanche recognition algorithms only makes sense if appropriate detection systems are available. The current state of our research enables us to contemplate, in the short term, to install a pilot system for the automatic detection of avalanches over areas extending up to several square km. The operation of such a pilot system will allow a significant improvement in the forecasting of risk.

Fig. 15

Fig. 16

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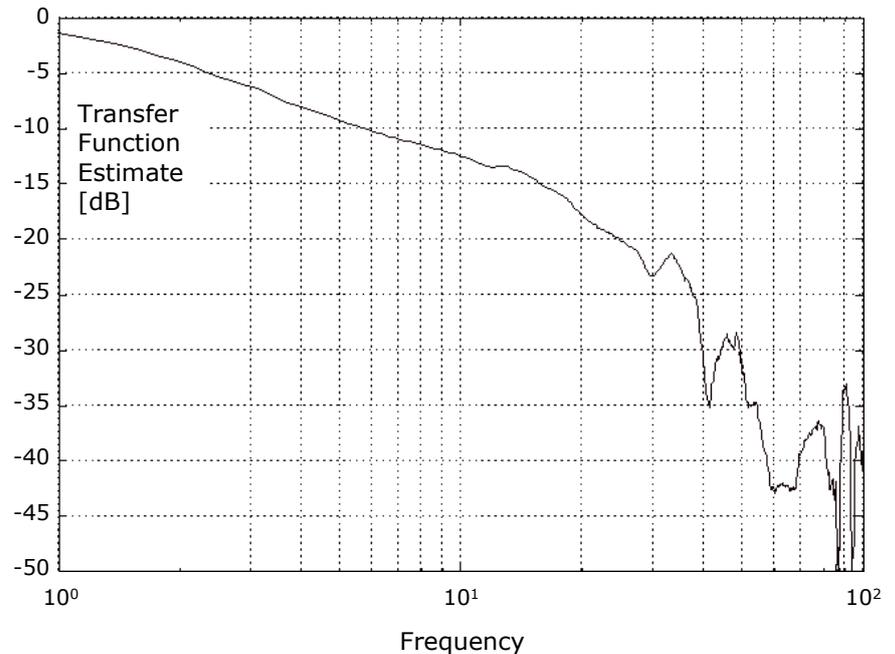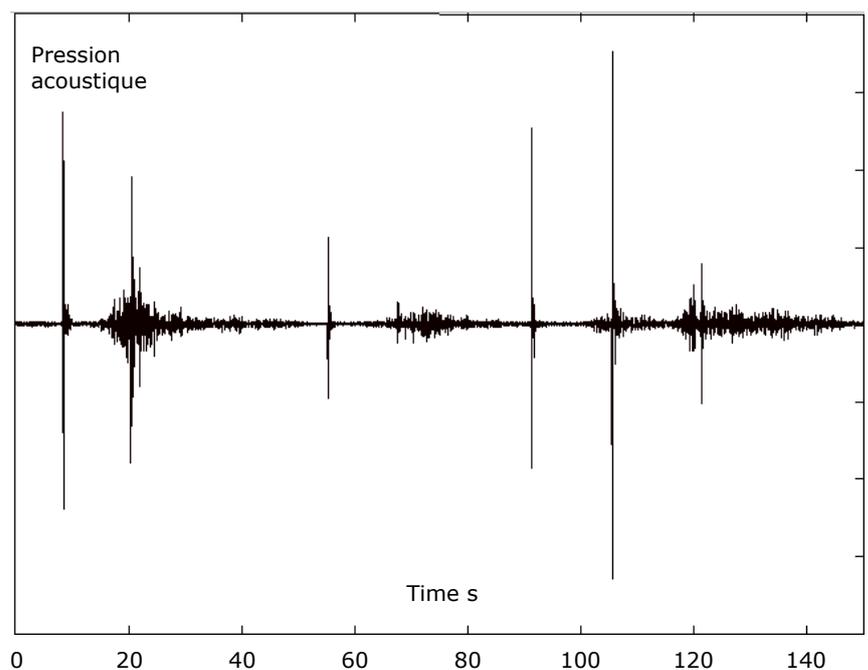Sommerfeld , R.A and H. Gubler 1977. Snow avalanches and acoustic emissions, *Annals of Glaciology*, **4**(83), 271-276.